



Evaluation of subpixel spectral methods for identification and change detection of secondary iron minerals

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Article Info

Abstract

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The main objective of this research is to discriminate secondary iron minerals and investigate changes caused by their relocation during mining activities using multispectral satellite data collected over an 11-year period. To achieve this aim, MF and CEM performed using USGS library spectra, field spectra and image spectra to identify secondary iron minerals and then the best result of the mineral identification stage implemented for changes detection of secondary iron minerals. The results of secondary iron minerals identification by MF method and images spectra have compatibility with the field data and laboratory analysis, so that the goethite is the most abundant secondary iron mineral in Darrehzar mine area. The change detection algorithm at the study area showed that mining activities and geochemical conditions cause change in the surface by transferring secondary iron minerals. The employed method including identifying minerals and detecting change of them allows users to find an effective technique for identifying the target material and apply it in the subtraction algorithm for change detection. Applying this method can be suggestions for future research in target change detection.

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1. Introduction

Mining activities cause changes in the Earth's surface by transferring rocks and surficial minerals such as ferruginous minerals in gossans (mostly hematite, goethite, and jarosite) and as a result of it, tailing damp and mine waste are produced (Atapour and Aftabi, 2007; Petropoulos et al., 2012; Vali et al., 2014) that could be detected and monitored using remote sensing instruments and techniques. Remote sensing instruments and images are usable for identifying reserves of ores, monitoring environment of mines, surveying tailings and identifying acid mine drainage (Paterson, 1997). Landsat series of satellites which have served remote sensing communities since the early 1970s, are helpful in monitoring the Earth's surface. Landsat 8 was launched in 2013, provides data for the VNIR, and shortwave infrared (SWIR) spectral bands with a 30 m spatial resolution as well as a 15 meter for panchromatic band in Operational Land Imager (OLI) sensor (Department of the Interior U.S. Geological Survey, 2016). Advanced Land Imager (ALI) sensor that was built to provide essential data for Landsat 8 mission, was successfully launched from Vandenberg Air Force Station, California on Earth observation (EO-1) satellite on November 21, 2000 (Beck, 2003; Lencioni et al., 2005). This sensor, which is very similar with Landsat especially in the visible portion, contains nine multispectral bands with 30 m spatial resolution (in VNIR and SWIR portions) (Mendenhall et al., 1998).

Various spectral techniques, such as Matched Filtering (MF), Constrained Energy Minimization (CEM) have been used to detect clay minerals bearing zones and Fe-minerals bearing area in order to map mineralized altered areas by advanced spaceborne thermal emission and reflection radiometer (ASTER) and Landsat multispectral images (Rowan et al, 2003; Zhang et

al, 2007; Gabr et al, 2010; Ghulam et al, 2010; Sadeghi, et al, 2013; Alimohammadi et al, 2015; Beiranvand Pour et al, 2015; Mokhtari and Seifi, 2021; Seifi et al., 2021). Many researchers have been involved implementing statistical (principal component analysis (PCA)) and spectral (mixture tuned matched filtering (MTMF) and linear spectral unmixing (LSU)) on ASTER, Landsat and Hyperion data for mapping ferruginous caps (gossans) and alterations zones in Sarcheshmeh and Darrehzar mines (Ranjbar et al, 2004; Beiranvand Pour et al, 2013; Hosseinjani Zadeh et al, 2014a; Hosseinjani Zadeh et al, 2014b; Abedi et al, 2015).

Many researchers have been involved in determining changes in mining areas in the Jharia coalfield (India), the central Appalachians, the Rutenburg mining region, Amazon, Ghana, the Greek island of Milos, Turkey and Germany using time series of Landsat, IRS, Quickbird and Worldview images (Prakash and Gupta, 1998; Ololade et al., 2008; Townsend et al., 2009; Schroeter and Gläßer, 2011; Petropoulos et al., 2012; Lobo et al., 2014; Yucel et al., 2014; Basommi et al., 2015). However these multispectral images are mainly used for mapping land use and land cover in mining areas (Prakash and Gupta, 1998; Ololade et al., 2008; Basommi et al., 2015), less attention is paid for detection of changes in minerals exposure for monitoring tailings. On the other hand, even though different algorithms were used for change detection, rare publications are available for mapping minerals through known spectral based algorithm such as spectral angle mapper (SAM) and matched filtering (MF) (Kruse et al., 1993; Harsanyi and Chang, 1994; Boardman et al., 1995; Wen and Yang, 2009; Valizadeh Kamran and Khorrami, 2018). However, many researchers reported the importance of MF and CEM in mineral identification (Rowan et al., 2003; Zhang et al.,

2007; Gabr et al., 2010; Sadeghi et al., 2013; Alimohammadi et al., 2015; Seifi et al., 2016; Mokhtari and Seifi, 2021; Seifi et al., 2021), its importance for discrimination of change detection is paid less attention.

Although many remote sensing studies conducted in the Darrehzar porphyry copper mine, in Kerman copper belt (KCB), south east of Iran (Ranjbar et al., 2004; Beiranvand Pour et al., 2013; Shahriari et al., 2013; Hosseinjani Zadeh et al., 2014a; Hosseinjani Zadeh et al., 2014b; Abedi et al., 2015; Seifi et al., 2016; Seifi et al., 2017; Seifi et al., 2019a; Seifi et al., 2019b), there are no investigations focuses on changes created by ore materials relocation using satellite images in this area. Hence the main objective of this research is to investigate changes created due to mining activities and discriminate iron minerals variation such as hematite, jarosite and goethite through the employed MF and CEM subpixel spectral techniques using multispectral satellite data at the Darrehzar mine during a 11-year period. For this purpose, MF and CEM are performing using United States Geological Survey (USGS) library spectra, field spectra and image spectra and then change detection of secondary iron minerals is going to implement according to the best result of the following stage.

2. Study Area

The Darrehzar porphyry copper deposit is located 80 km northeast of Sirjan in Kerman province and eight km south of well-known Sarcheshmeh porphyry copper deposit. This area is situated in the Urumieh–Dokhtar magmatic belt of Iran (Sahand- Bazman) and is a part of the Kerman copper belt, namely Dehaj-Sarduiyeh magmatic arc (Figure 1).

In the years 1969 and 1970, a former Yugoslavian geological team conducted an initial exploration investigation in the Darrehzar region on behalf of the Iranian Geological Survey. They carried out geological, geophysical, and geochemical investigations before drilling 24 wells and two tunnels for exploration. On the basis of these facts, they calculated the Cu/Mo reserve. Additional drilling and new exploration activities started in 1993, and overburden removal efforts began in 1995. In order to produce the necessary ore for semi-industrial bioheap leaching studies, extraction operations have been carried out since May 2004 for a period of one year. Due to the presence of sulfide-bearing minerals, water infiltration, and oxidant bacteria at the deposit site, acid mine drainage was produced (Moazzen zadeh et al., 2006). Official exploration efforts for ore reserve estimation and mine design began in 2006, and an operational license was then secured from the Kerman province's Industries and Mines Organization in 2007 (Yazdanpanah, 2011; Hosseinjani Zadeh, 2013; Parsapoor, 2014).

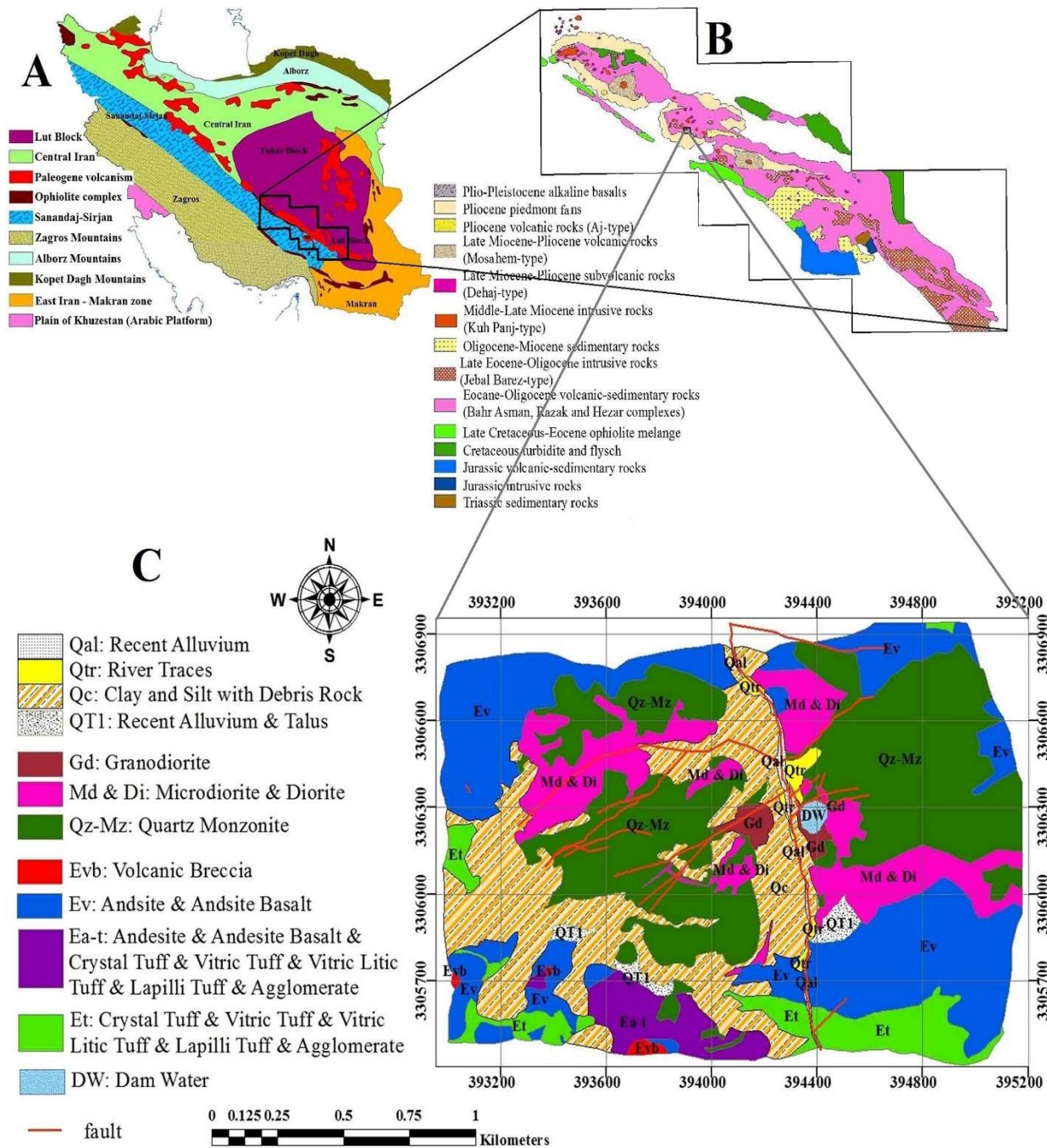


Figure 1. A) A structural-sedimentary map of Iran (modified after Moeen Vaziri, 2004), B) geological map of Deharsarduyeh (modified after Nateghi and Hezarkhani, 2013), and C) geological map of the Darrehzar porphyry copper deposit (modified after Alizadeh Sevari and Hezarkhani, 2014).

The Darrehzar porphyry is situated in a diorite diorite–quartz pluton of Oligocene–Miocene age which intrudes an Eocene Volcanic–Sedimentary complex consisted mainly of volcanoclastics, andesite, trachyandesite and sedimentary rocks. Hydrothermally altered rocks are highly fractured, and supergene alteration has produced extensive limonite and leaching of sulfide, giving a characteristic reddish or yellowish color to the altered rocks (Geological Survey of Iran, 1973). Propylitic and phyllic alterations are dominant in surface rocks with sporadic small areas of argillic alteration. Propylitic alteration is well developed in the area and extends a few hundred meters around the central phyllic zone. The phyllic alteration persists below the oxidation zone and is characterized by enrichments in Si, Al and K, and depletions in Na, Ca, Mg and Mn (Ranjbar et al., 2001). Potassic alteration is not observed on surface, possibly as a result of a surface related weathering or intense phyllic overprint (Geological Survey of Iran, 1973).

3. Materials and Methods

3.1. Satellite data

In order to investigate the mineral identification and change detection in Darrehzar area, images were selected in a manner to cover two periods in summer (normally AMD observed in summer) (Guo et al., 2013; Navarro Torres et

al., 2011); a) 2004 (before official exploration activities (2005) and after reported acid mine drainage pollution (2003) (Hosseinjani Zadeh, 2013; Moazzen zadeh et al., 2006)) and b) 2015 (before the first implemented field investigation by authors at the study area). The image of 2015 was related to OLI sensor of Landsat 8 which has been acquired on the 9 August 2015 and was downloaded from the USGS website. Unfortunately, there was no suitable Landsat (ETM + and TM) images in 2004 for the study area. Therefore, the ALI image, which is similar to Landsat in terms of sensor design, platforms orbital characteristics, spatial and spectral resolution, was chosen (Department of the Interior U.S. Geological Survey, 2016). The ALI image, which is contemporary in season with the OLI, was acquired on the 26 July 2004 from USGS website.

3.2. Field data

Field reconnaissance was conducted for collecting rock samples at the study area in 2016. Fresh/weathered and representative oxidized rocks samples were collected from different parts of mining pit and tailing dumps through a random punctual sampling and their positioning were recorded by global position system (GPS) Oregon 650. These samples were collected from different areas of the mining pit and waste dumps of the Darrehzar mine (Figure 2).

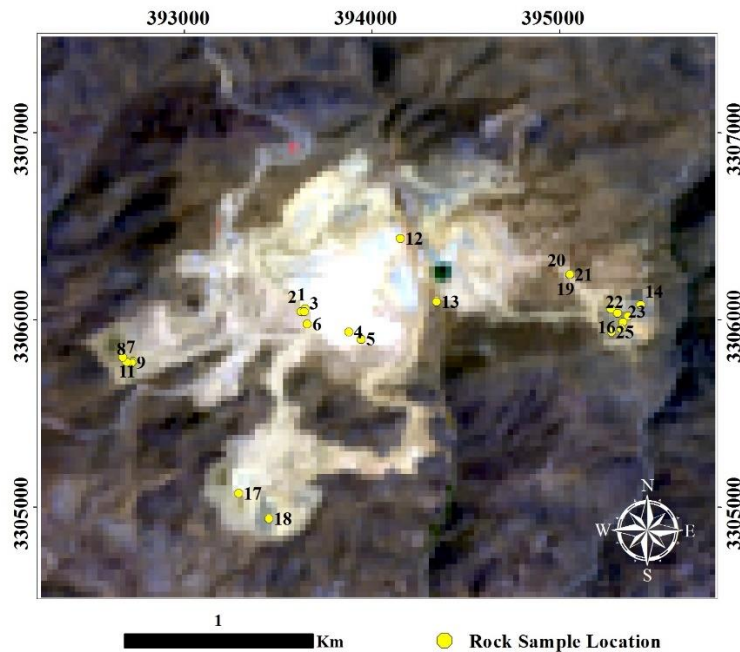


Figure 2. Location of rock samples which were plotted on the color composite of OLI image (each number on the figure indicate sample number such as; 1: Da1, 2: Da2, 3: Da3, 4: Da4, 5: Da5, 6: Da6, 7: Da7, 8: Da8, 9: Da9, 10: Da10, 11: Da12, 12: Da21, 13: Da24, 14: Da26, 15: Da27, 16: Da28, 17: SD, 18: SD1, 19: NED, 20: NED1, 21: NED2, 22: ED1, 23: ED2, 24: ED3, 25: ED4).

Collected samples were sealed in clean polythene bags and kept for further analysis (spectroscopy studies and X-ray diffraction). Spectroscopy of rock samples were carried out in a laboratory by an analytical spectral device (ASD) FieldSpec®3 spectroradiometer as well as spectral processing and mineral recognition were conducted using PIMA View V.3.1., Specmin-PRO V.3.1, and ENVI 5.3® software. The average percentage of minerals in results of PIMA View and Specmin-PRO for the spectra of each sample was considered as the final percentages for minerals in samples. The XRD spectra of powdered samples with the best spectral features were obtained with a Philips PW 1800 diffractometer using Cu-K α radiation ($K\alpha = 1.542 \text{ \AA}$), operating at 40 kV and beam current of 30 mA. Moreover, previous researches

(Abdolzadeh, 2005; Atapour and Aftabi, 2007; Keshavarzi, 2006) in the study area were used to validate results of image processing of ALI data (2004).

3.3. Pre-processing:

Applying atmospheric correction algorithms is a crucial step in the pre-processing process because atmospheric water vapor and particles cause noise in the digital values of the pixels in remotely sensed images (Hosseinjanizadeh, 2013). Therefore, the ALI and OLI data were corrected using the Fast Line-of-sight Atmospheric Analysis of Spectral Hypercubes (FLAASH) method. The implemented atmospheric correction removes atmospheric effects on image pixels and retrieves spectral reflectance of Earth's surface from multispectral radiance images (ITT Visual

Information Solutions, 2008). Level 1T of ALI and OLI images were downloaded from the USGS website and were geometrically accurate by 1:25000 topographic map and stream layer of study area.

3.4. Endmember extraction:

In order to identify secondary iron minerals including; hematite, goethite and jarosite, and detect changes for them in aforementioned period,

the MF and CEM algorithms were implemented on ALI and OLI images separately using spectra of these minerals from USGS spectral library. Whereas there are many spectra for each mineral in the USGS spectral library hematite2, goethite3, and jarosite1 spectra were chosen according to the research of Hosseinjanizadeh et al. (2014b) (Figure 3A).

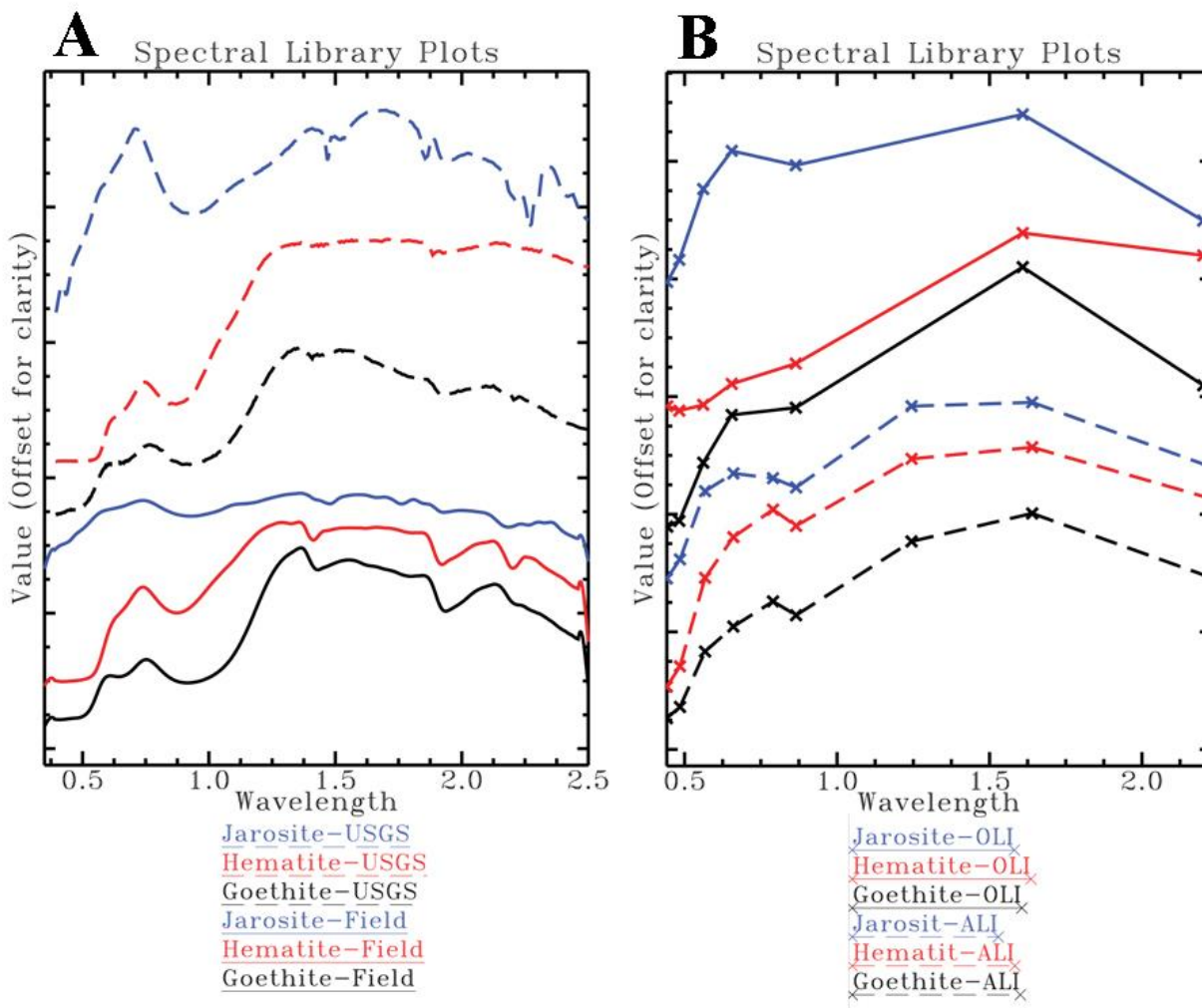


Figure 3. The USGS, field and image spectra of secondary iron minerals such as jarosite, hematite and goethite: A) The USGS and field spectra; B) The OLI and ALI image spectra.

As mentioned above, minerals in rocks samples were identified by spectroscopy measurement

using an ASD FieldSpec®3 spectroradiometer in a laboratory. So that samples are mixtures of

different minerals and less abundant minerals may not be observed in one single measurement, multiple spectra were acquired at the laboratory from rocks samples taking into account (95 spectra from 26 samples). The spectral processing and mineral recognition were carried out by PIMA View V.3.1., Specmin-PRO V.3.1, and ENVI 5.3® software.

Goethite, hematite and jarosite in rock samples are identified using absorption features in ASD spectra because of crystal field transitions of ferric/ferrous iron in Fe-minerals, OH stretch in secondary iron minerals and Fe-OH stretch in jarosite (Zabcic 2008; Hosseinjanizadeh et al. 2014b; Seif et al. 2017). At the present study, absorption features of crystal field transitions in ferric/ferrous iron (0.62-0.63 and 0.90-0.93 μm in goethite, and 0.87 μm in hematite) and OH stretch (1.41 and 1.92 μm in hematite, and 1.42 and 1.93 μm in goethite) combinations were used for identifying of goethite and hematite in rock samples. Absorption features associated with crystal field transitions of ferric/ferrous iron (0.44 μm in jarosite) and Fe-OH (2.26 μm in jarosite) were considered as jarosite (Figure 3A). Moreover, Al-OH absorption feature (2.20 μm) observed in ASD spectra could be because of the presence of clay minerals in measured samples (Hosseinjanizadeh et al., 2014b, Seifi et al., 2019a).

For pure spectra extraction of images' pixels, at first, inherent dimensionality of image data was determined using Minimum Noise Fraction (MNF) rotation to separate noise in the data, and to decrease the computational requirements for next processing. Then Pixel Purity Index (PPI) were performed to segregate the purest pixels in the Landsat 8 and ALI multispectral images (Research Systems Inc., 2004). Corresponding spectra of bright pixels in the PPI image and their similarity with resampled spectra of spectral

library were used for extracting the purest spectra. Moreover, Sequential Maximum Angle Convex Cone (SMACC) spectral tool was carried out to find spectral image endmembers and to help the pure spectra extraction process (Research Systems Inc., 2004).

There are absorption features associated with crystal field transitions of ferric/ferrous iron (0.43-0.48, 0.50-0.67 and 0.85-0.94 μm) in band 2 ALI and OLI for jarosite, in band 3 OLI for hematite, and in band 5 OLI and band 6 ALI for jarosite, goethite and hematite. As well as, ALI and OLI spectra presented absorption characteristics of OH stretch (1.9 μm) in band 9 ALI and band 7 OLI for three mentioned minerals (Figure 3B).

3.5. Mineral Identification:

Goethite was chosen for selecting the best method because it is the most stable Fe-mineral in acid mine drainage participates and gossans in mining area (Atapour and Aftabi, 2007; Singh et al., 1999). In the present research, CEM and MF subpixel spectral techniques were implemented using goethite spectrum of USGS spectral library, OLI image goethite spectrum, and field goethite spectrum on Landsat 8 (OLI) data. MF technique is a procedure implements a partial spectra unmixing to evaluate the user-defined endmembers abundance from a set of reference spectra for mapping the surface minerals. In fact, this method minimizes the response of the composite unknown background by projecting vector of each pixel onto a subspace that is orthogonal to the background spectra and then maximizes the response of the known end-member by comparing the residual pixels to each of the reference spectra and finally matches the known signature (Harsanyi et al, 1994; Boardman et al, 1995; Rowan et al, 2003). CEM technique is a performance of matched filtering methods in the hyperspectral analysis context. CEM applies a

finite impulse response filter (a linear operator) to pass through the desired target while minimizing output energy of background. A correlation or covariance matrix is used for characterization of the component unknown background and increased the contrast between the target and background spectra (Farrand and Harsanyi, 1994; Farrand and Harsanyi, 1997; Resmini et al, 1997; Settle, 2002; ENVI 5.3 online help). Digital Numbers (DNs) of CEM and MF image pixels present relative percentage of interest materials per pixel, as a result they were classified in three groups, 5% - 20% (low value), 20% - 50% (medium value), and 50% - 100% (high value) according to ASD and XRD data in the research of Seifi et al. (2019a). The best results of CEM and MF subpixel spectral techniques corresponding the ASD data was chosen as the best method for next stage, mineral identification. The best method was used to identify secondary iron minerals containing, hematite, goethite, and jarosite by ALI and OLI data. The results of secondary iron minerals identification were classified also in three groups, 5% - 20% (low value), 20% - 50% (medium value), and 50% - 100% (high value). The results of OLI image validated by ASD FieldSpec®3 spectroradiometer data (Table 1). The results of ALI image validated by information in previous researches (Abdolzadeh, 2005; Atapour and Aftabi, 2007; Keshavarzi, 2006) in study area.

3.6. Change Detection

The employed method for change detection of secondary iron minerals which is practical to identify change of the target material includes three phases, method selection, target identification and change detection. The method allows users to find an effective technique for detecting the target material and apply it in the subtraction algorithm for change detection. The

gray-scale MF images of following stage were used for detecting hematite, goethite and jarosite changes in an 11-year period between 2004 (ALI data) and 2015 (OLI data). For this purpose, differences between MF images of secondary Fe-minerals are computed by subtracting the MF image of 2004 from MF image of 2015 and classes are determined by applying the certain value of statistical threshold. In order to select threshold in subtraction algorithm, a no change class should be surrounded by an equal number of positive and negative classes between -1 and +1 (Research System Inc., 2004). So that, the same values of statistical data (0.2 and 0.5) for classifying secondary iron minerals images in following stage were used as statistical threshold.

4. Results and Discussion

4.1. Method Selection

Iron minerals changes created by mining activities investigated through the employed MF and CEM subpixel spectral techniques using multispectral satellite data at the Darrehzar mine during a 11-year period. Goethite was used to select the best method for next stage, mineral identification, because it is one of the most stable secondary iron minerals in a wide range of pH (Seifi et al., 2019a; Singh et al., 1999). Goethite was detected using MF and CEM methods with the help of USGS library, image and Field spectra to choose the best method for identifying secondary iron minerals. The results of goethite identification were classified into three categories, 5%-20% (low value), 20%-50% (medium value), and 50%-100% (high value). Goethite is only found in the mining area, according to the results of the MF method's detection by the USGS spectral library and Field spectra, therefore only 5 to 20% of the desired pixels contain goethite. (Figure 4A and C). Additionally, the results of

goethite identification using the CEM method by the USGS spectrum library and Field spectra show that this mineral is found in the mining area with a medium value (20 to 50%) and in the vicinity of the mine with a low value (5 to 20%). (Figure 4D and F). In the results of detecting goethite using MF and CEM methods by image spectra, this mineral is identified in the mining area and also around the mine with low (5 to 20%), medium (20 to 50%) and high values (50 to 100%). (Figure 4B and E). In the results of the MF method by the OLI image spectrum are presented that there is goethite in the mining area with low, medium and high values, so that the highest value of goethite is

located in the mine pit, and the tailings dumps show medium and low values (Figure 4B). The results of the CEM method by the OLI image spectrum reveal the low, medium and high values of goethite, so that there is the high value of goethite in the mining pit, and the medium and low values of goethite are situated in a wide area around the mine (Figure 4E). Among the results of the six methods, the result of the MF method by image spectra has compatibility with the field data and laboratory analysis, so that the goethite mineral is observed in most of the collected samples (Figure 2 and Table 1).

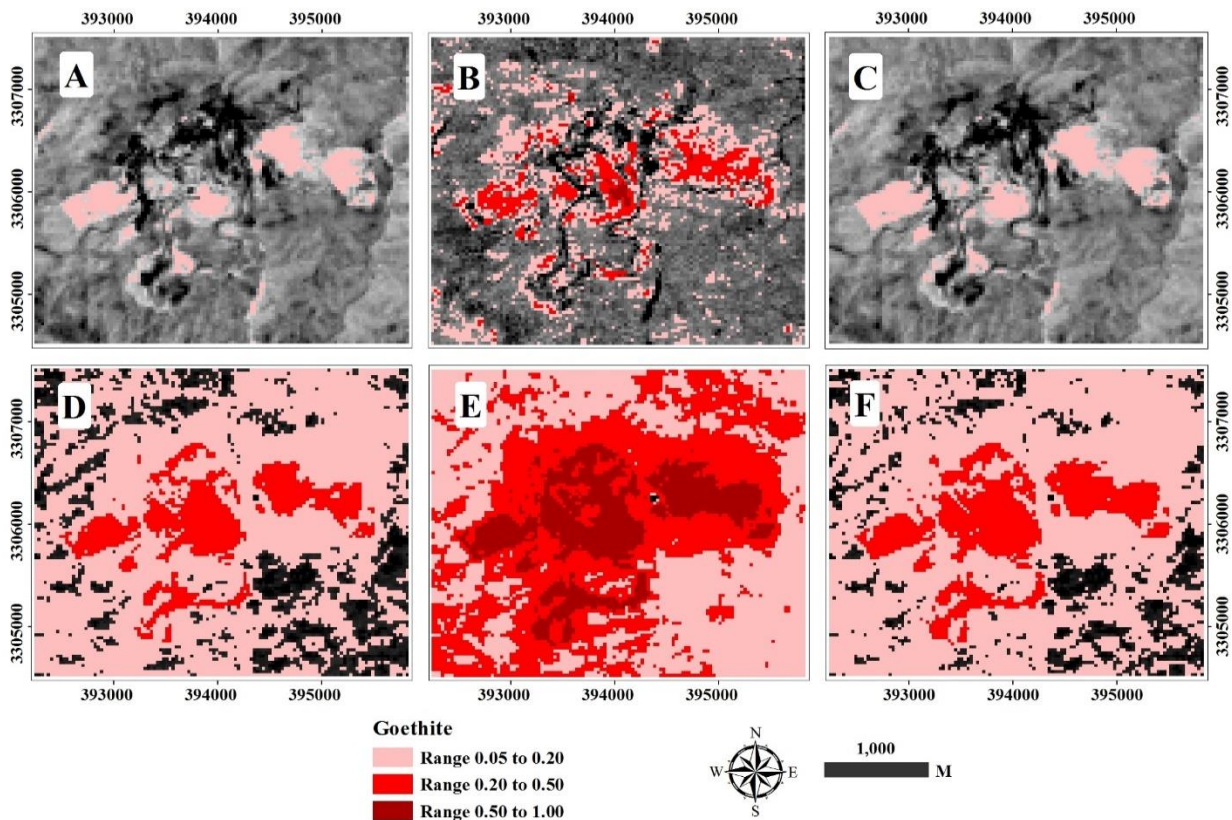


Figure 4. Spatial distribution maps of Goethite in the study area created by A) the MF method and USGS spectrum; B) the MF method and OLI image spectrum; C) the MF method and field spectrum; D) the CEM method and USGS spectrum; E) the CEM method and OLI image spectrum; F) the CEM method and field spectrum.

4.2. Mineral Identification

Secondary iron minerals containing goethite, hematite and jarosite were identified by MF method using images spectra. The spectra of aforementioned minerals were extracted from ALI and OLI satellite images (Figure 3B) and used for detecting secondary iron minerals. Goethite is the most abundant secondary iron mineral in Darrehzar mine area, and it presents a few changes between 2004 to 2015 (Figure 5A and D). Hematite is mainly detected with moderate (20-50%) and high (50-100%) values along the Darrehzar River, which intersects the Darrehzar mine into two parts, in 2004 image (ALI) (Figure 5B). But it is identified by low value (5 to 20%) in

the mine area and also by medium value (20 to 50%) in some central parts of the mine in the OLI image (2015) (Figure 5E). Jarosite is detected using the ALI satellite image (2004) in the mining area with low (5 to 20%), medium (20 to 50%) and high (50 to 100%) values, so that the low amounts of jarosite is detected in a wider area than the medium and high values in Darrehzar mine (Figure 5C). Moreover, this mineral is identified with medium and high values in a more widespread area than the low value of jarosite in the Darrehzar mine in the 2015 image (OLI), and mostly these values correspond to the pit of Darrehzar mine (Figure 5F).

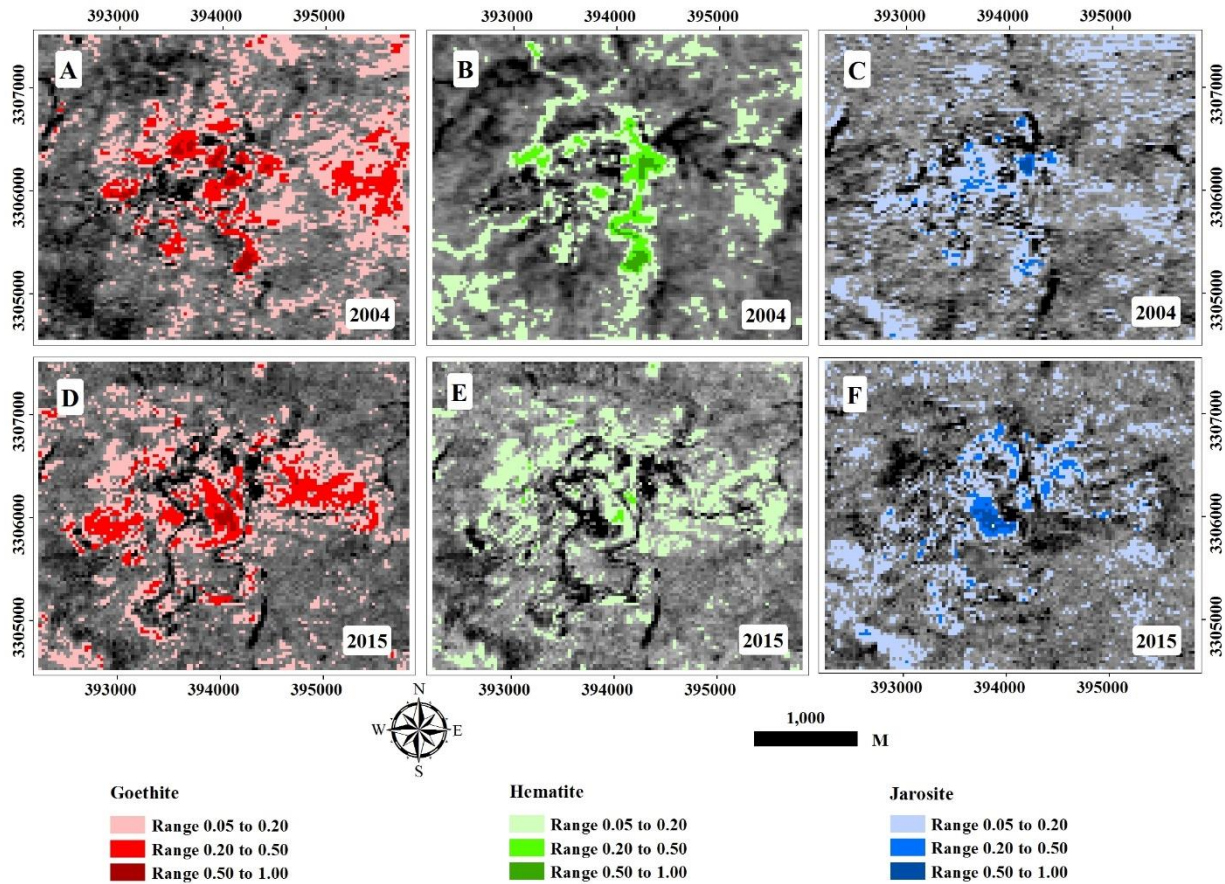


Figure 5. A) Result of the MF method and ALI image spectrum for goethite identification; B) Result of the MF method and ALI image spectrum for hematite identification; C) Result of the MF method and ALI image spectrum for jarosite identification; D) Result of the MF method and OLI image spectrum for goethite identification; E) Result of the MF method and OLI image spectrum for hematite identification; F) Result of the MF method and OLI image spectrum for jarosite identification.

Field observations and laboratory analysis of collected samples from the mine confirm the image processing results.

Previous studies approve the results of ALI images processing so that Abdolzadeh (2005) reported that there are jarosite, goethite, hematite and limonite in oxidation zone of Darrehzar mine as well as Atapour and Aftabi (2007) mentioned that there is a range of goethite > hematite > jarosite in gossan and leached zone of Darrehzar. Furthermore, ICP-MS results on by Keshavarzi (2006) revealed high amounts of iron (Fe) in collected sediments of Darrehzar River.

The spectral analysis of collected samples spectra (taken from Darrehzar mine in 2015) also present the following results:

- Goethite is the most abundant mineral present in the collected samples so that there is it in most of the collected samples with high values (50 to 100%).
- The samples taken from the mining pit show medium (20 to 50%) and high (50 to 100%) values of hematite, while collected samples from the tailing dumps present a low value (5 to 20%) of that.

- Jarosite is presented in the samples taken from Darrehzar mine area with low value (5 to 20%) (Figure 2 and Table 1).

The X-ray diffraction from rock samples correspond to the location of discriminated areas

showed existence of jarosite and goethite specially inside the Darrehzar mine. These minerals are observed mainly with clay minerals in XRD analysis from rock samples (Seifi et al., 2019a).

Table 1: Characteristics of rock samples measured by ASD FieldSpec®3 spectroradiometer (The final percentage for minerals in the table were calculated by average percentage of minerals in results of PIMA View and Specmin-PRO).

Mining Pit										
	Da1	Da2	Da3	Da4	Da5	Da6	Da21	Da24		
Goethite	56.08	0	47.43	62.72	58.09	27.37	41.83	26.03		
Hematite	0.59	0	0	0	41.91	35.43	58.17	50.92		
Jarosite	0	0	0	12.79	0	6.37	0	0		
Clay	43.33	100	52.57	24.49	0	30.83	0	23.05		
Western Dump										
	Da7	Da8	Da9	Da10	Da12		Southern Dump			
							SD	SD1		
Goethite	80.31	84.66	77.37	39.51	0		Goethite	77.81	82.1	
Hematite	6.55	9.46	0	12.01	70.74		Hematite	3.16	2.66	
Jarosite	0	0	0	0	0		Jarosite	6.16	0.48	
Clay	13.14	5.88	22.63	48.48	29.26		Clay	12.87	14.76	
Eastern Dump										
	Da26	Da27	Da28	ED1	ED2	ED3	ED4	NED	NED1	NED2
Goethite	72.07	65.54	75.69	72.75	59.39	78.91	61.75	53.52	68.29	51.18
Hematite	16.94	0	3.38	4.17	0	21.09	0	7.66	7.96	4.31
Jarosite	10.99	0	9.52	0	0	0	0	23.77	14.22	29.04
Clay	0	34.46	11.41	23.08	40.61	0	38.25	15.05	9.53	15.47

4.3. Change Detection

Subtraction Algorithm was implemented on the results of the previous step, mineral identification, in order to detect the changes of secondary iron minerals such as goethite, hematite and jarosite, as well as statistical threshold (-0.5, -0.2, 0, 0.2, 0.5) was used to achieve better results. The result of change detection is based on the brightness of pixels in the initial and final states of images as ascending and descending of secondary iron minerals (jarosite, goethite, and hematite) in the mine site during an 11-year period in red and in blue, respectively. Implementation of change

detection algorithm at the area revealed that mining activities and geochemical conditions cause change in the surface by transferring minerals such as hematite, goethite, and jarosite. Goethite presents both increase and decrease in different areas of the mining pit and western, eastern, and southern tailing dumps, so that it shows the displacement of the goethite in the mine (Figure 6A). Hematite mainly shows a decrease in the area of mine but increases sporadically in the mining pit and the paths leading to the waste dumps (Figure 6B). Jarosite increases mostly in mining pit and in small parts of western, eastern

and southern waste dumps. The Darrehzar river is the most important place where jarosite is increased because the waste water reacts with pyrite and minerals such as jarosite are deposited. It should be mentioned that the surface water lake where the mine's acidic waters enter, is situated in the Darrehzar mine (Figure 6C). According to the history of the mine, the inactive western dump has been created and expanded during 1993 to 2010.

This dump which is the oldest actually coincides with the exploration operations in 1993 (Yazdanpanah, 2011). The inactive eastern dump has been generated during 2010 to 2015. Furthermore, the active southern waste dump created in 2010 (unpublished data from the Darrehzar mine).

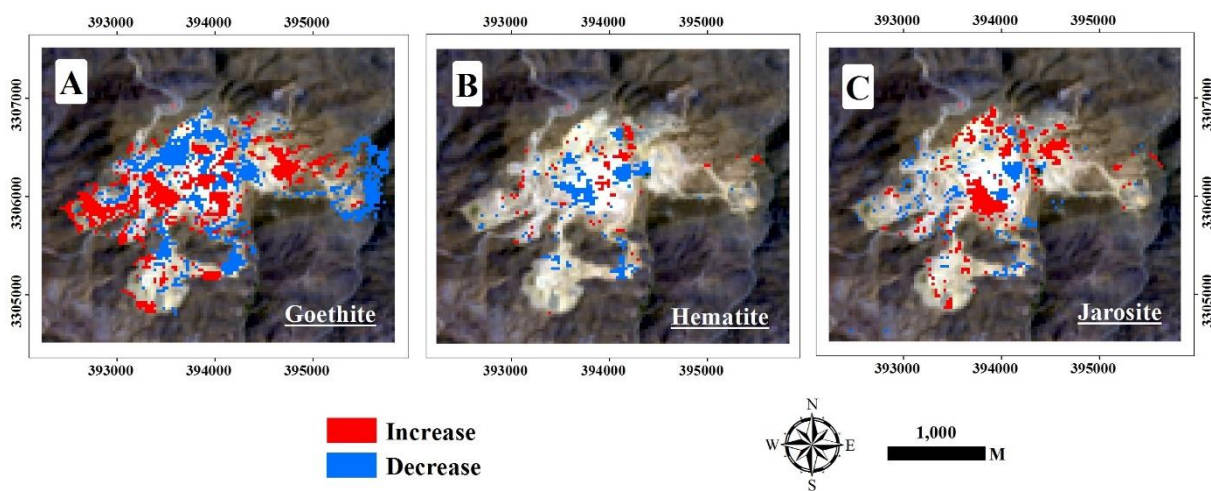


Figure 6. A) Result of the goethite change detection; B) Result of the hematite change detection; C) Result of the jarosite change detection. (Blue: decrease in goethite, hematite and jarosite; Red: increase in goethite, hematite and jarosite).

5. Conclusion

The employed method is practical to identify change of the secondary iron minerals includes three phases, method selection, mineral identification and change detection. Sub pixel methods determine approximate values of minerals in image's pixels. Among the results of

the six methods, MF and CEM using USGS library, image and Field spectra, the result of the MF method and image spectra is the best method for secondary iron minerals identification. The results of secondary iron minerals identification by MF method using images spectra have compatibility with the field data and laboratory analysis, so that the goethite is the most abundant secondary iron mineral in Darehazar mine area. Identifying target and detecting change of it allow users to find an effective technique for identifying the target material and apply it in the subtraction algorithm for change detection. The subtraction algorithm was performed on the results of the

mineral identification stage for changes detection of secondary iron minerals and statistical threshold was used. Implementation of change detection algorithm at the area revealed that mining activities and geochemical conditions cause change in the surface by transferring minerals such as hematite, goethite, and jarosite. Applying this method can be suggestions for future research in material change detection.

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